Università degli Studi di Bologna Dottorato di Ricerca in Geofisica – XXIII Ciclo

# MODELLI DINAMICI DI ROTTURA SISMICA

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#### **1.** EARTHQUAKE SOURCE DYNAMICS

- Elasto dynamic problem
- Rupture description
- Dislocation vs. crack models
- Forward modeling scheme
- Rupture stages

#### 2. FAULT GOVERNING LAWS ( CONSTITUTIVE EQUATIONS )

- Fault models
- Physical phenomena in faulting
- Fracture criteria and constitutive laws
- Strength and constitutve laws
- Slip dependend friction laws
- Rate and state dependent friction laws

#### 3. EARTHQUAKE NUCLEATION

#### **4.** RUPTURE PROPAGATION IN 2 – D FAULT MODELS

- Numerical methods
- BIE vs. FD
- Slip weakening vs. Dieterich Ruina law
- The cohesive zone and the breakdown processes
- Theoretical interpretations and correspondency formula
- The estimate of  $d_0$  and related problems
- The importance of the evolution equation

#### 5. RUPTURE PROPAGATION IN A TRULY 3 – D FAULT MODEL

- The numerical method
- The reference case. Comparison between 2 D and 3 D models
- Coupling of two modes of propagation. The rake variation
- Dependence on the absolute stress level. Simmetry leak
- Heterogeneous configurations

#### 6. RHEOLOGICAL HETEROGENEITIES, ARREST AND HEALING PHENOMENA

CRACK

- The crack and arrest models
  - The barrier healing
  - The pseudo self healing
- The pulse
  - Different materials
  - Analytical modifications to friction laws
  - \_\_\_\_\_The Nielsen' s model

#### 7. CONVERGENCE

#### 8. FAULT INTERACTION AND STRESS TRIGGERING

- The spring slider (SS) model
- Analytical stress perturbations: step and pulse
- Realistic stress perturbation: a complex stressgramma
- Application to the June 2000 SISZ seismic sequence I: SS fault model
- Application to the June 2000 SISZ seismic sequence II: 3 – D fault model

# Earthquake source dynamics

# Elasto - dynamic problem

\* Solution of the fundamental elasto – dynamic equation ( i. e. the II law of dynamic for continuum media ):

$$\rho(d^2/dt^2)U_i = \sigma_{ij,j} + f_i$$
;  $i = 1, 2, 3$ 

where:

- $\rho$  is the mass cubic density,
- **U** is the particle displacement vector ( $\mathbf{U} = \mathbf{x}' \mathbf{x}$ ),
- { $\sigma_{ij}$ } is the stress tensor;  $\sigma_{ij} = C_{ijkl}e_{kl}$ ; i,j,k,l = 1, 2, 3, where  $C_{ijkl}$  is the elastic constant tensor, accounting for the rheology of the medium and  $e_{kl}$  is the strain tensor ( $e_{kl} = \frac{1}{2} (U_{k,l} + U_{l,k})$ ),
  - is the body force vector.

Choice of the climensionality d of the problem
 (1 – D, 2 – D, 3 – D).
 (d = rank of the U array, i. e. number of equations)

**1.** Wave propagation problem: Hyperbolic PDE D'Alembert wave equation:  $\nabla^2 U - (1/c_0) (\partial^2/\partial t^2) U = 0$ where  $c_0$  is the wave speed.

2. Rupture problem

# **Rupture** Description

Following Scholz ( 1990 ) the rupture can be described by using:

#### CRACK MODELS

The energy dissipation at crack edge (or crack tip) is paramount. Describe explicitely the crack propagation.

#### FRICTION MODELS

The effects at the edges are not explicitley considered. Explicitly allow for the calculation of the evolution of stress tensor components in terms of material properties of the fault.

# **Dislocation vs.** Crack Models

#### **DISLOCATION MODELS**

- \* Study of displacement discontinuity
- Slip is assumed to be constant on the fault;
   The fault evolution is represented by unilateral or bilateral motion (rectangular dislocations: Haskell's model)
- rupture front and it neglects dynamics of faulting
- **1** Long period seismic waves modeling (  $\lambda \ge L_{fault}$  )
- constant dislocation is inadmissible; strain energy at crack tip is unbounded; stress drop is infinite

#### **CRACK MODELS**

- \* Impose finite energy flow into the rupture
- \* Slip is not prescribed,

but it is calculated from the stres drop and from the fault strength S<sup>fault</sup>

crack ( after nucleation processes ), increases the stress outiside the crack near the crack tip ) and tends to facilitate further grow of the rupture

- The motion is determined by fracture criterion ( and eventually by the assumed constitutive law on the fault )
- The problem is characterized by assuming the boundary conditions on the fault plane. It has mixed b. c.: slip assigned outside the crack tip and stress tensor components inside the crack tip

# **Forward modeling scheme**

#### **1.** Fault model:

- Fault geometry (orientation, planar or non planar, ...)
  - Fault system (multiple segments, multiple faults, ... )

#### 2. Meclium surrounding the fault surface( s )

Properties of the medium surrounding the fault(s): cubic mass density structure, velocity structure, anysotropy attenuation

3. Choice of the dimensionality d' of the problem (1 – D, 2 – D, 3 – D, 4 – D).
(d' = number of the independent variables in the solutions)

4. Choice of the representation

**5.** Choice of the numerical method - (FE, FD, BE, BIE, SE, hybrid)

#### 6. Specification of the Boundary Conditions

- Domain Boundaries Conditions (DBCs)
- Fault Boundary Condition (FBCs)
- Auxiliary Conditions (ACs)

#### 7. Specification of the Initial Conditions

Initial conditions on the fault: (initial slip, slip velocity, state variable, pre – stress);

- Initial conditions outside the fault: (tectonic load, (state of neighbouring faults: the fault is <u>not</u> an isolated system))
- 8. Evaluation of the solutions

Convergence analysis ( consistency + stability )



#### 1. Nucleation (quasi – static to dynamic evolution)

- How can we simulate nucleation?
- How can we promote fault instability?

#### 2. Propagation

- What is the fault constitutive equation (governing law)?

#### 3. Healing

- What type of healing occurs?
- What controls fault healing?

#### 4. Rupture arrest

- What is responsible of rupture arrest?
- How can we represent it? Earthquake energy balance?

#### 5. Fault re – strengthening

- How can we model further instabilities episodes on the fault?

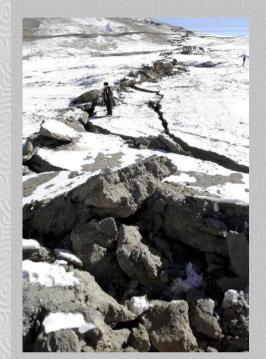
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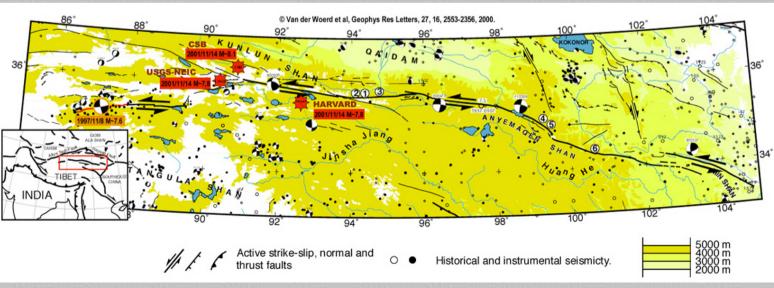
# Support Slides: Parameters, Notes, etc.

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# Geometrical complexity

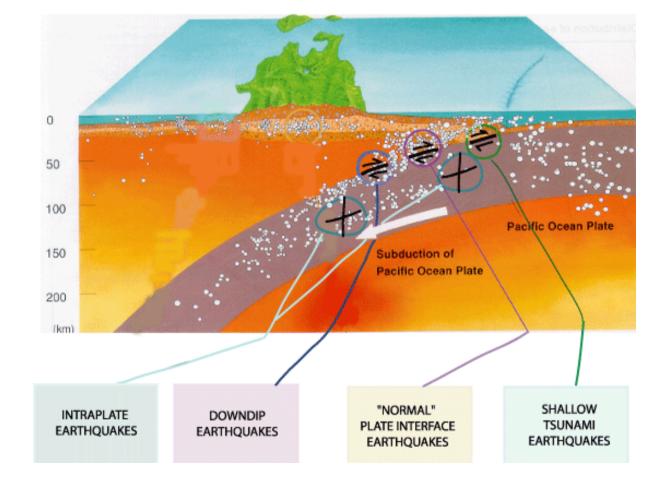


Kokoxili *M*<sub>w</sub>, 7.9 earthquake (Qinghai Province, China) Strike Slip Surface Breaks

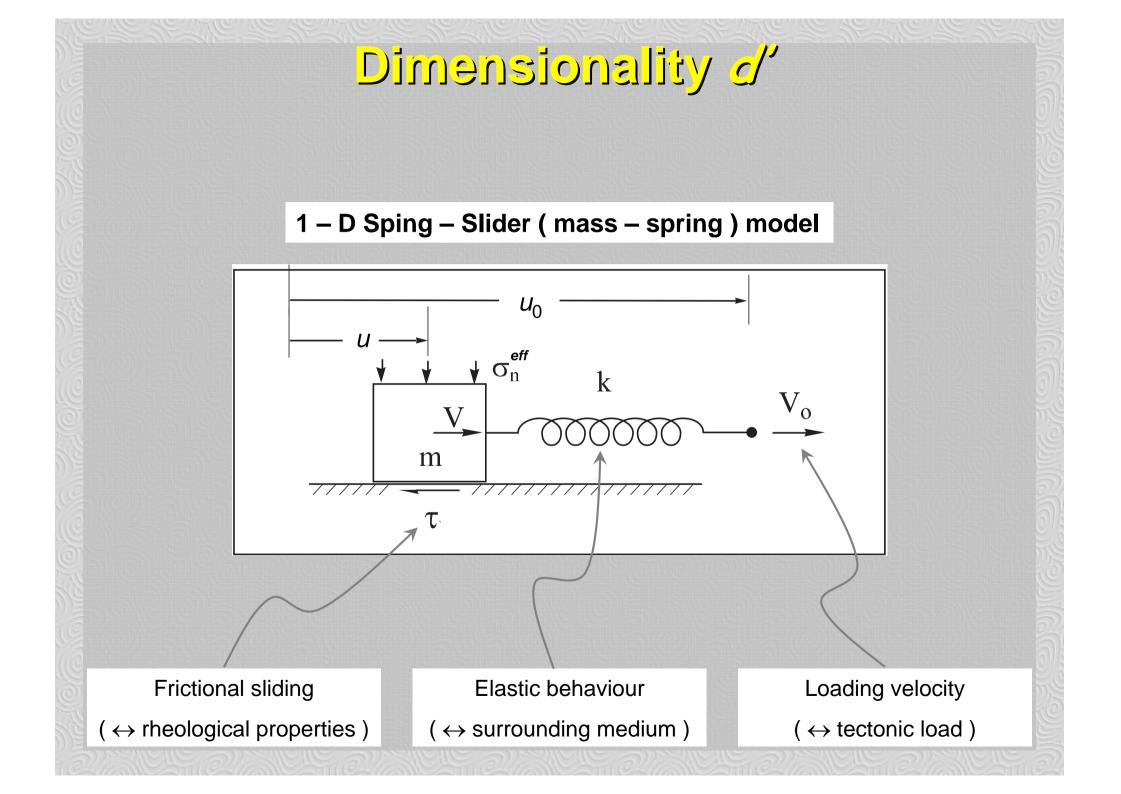


# Different types of earthquakes

Interplate
 Tsunami
 Crustal
 Downdip
 Intraplate
 Deep



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Let we recall basics concepts on dislocation theory.

If  $U_{i,j}$  is continuous in the integration domain  $\int_{P_1}^{P_2} U_{i,j} dx_j$  then does not depends on the integration path and therefore:

$$\oint_C \mathrm{d} U_i = 0, \qquad \forall C.$$

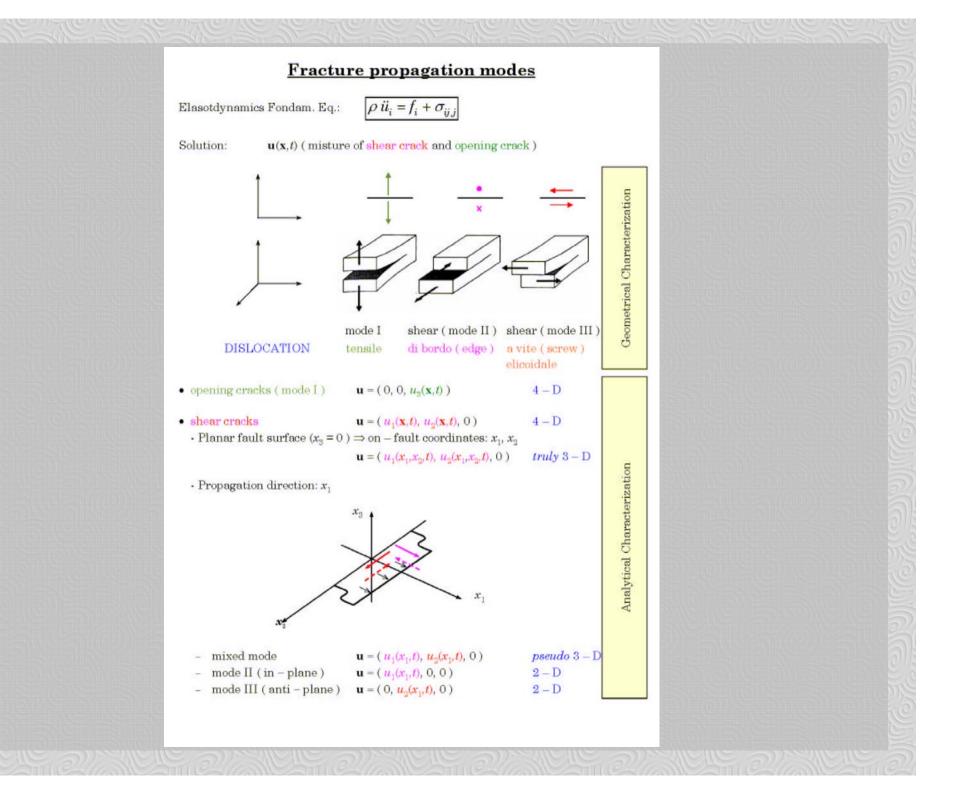
On the contrary, when  $\oint_{C} dU_i = b_i$  the considered body contains a **dislocation** and the circuit *C* contains at least one curve ( the **dislocation curve** ) on which the tensor  $U_{i,i}$  is not defined.

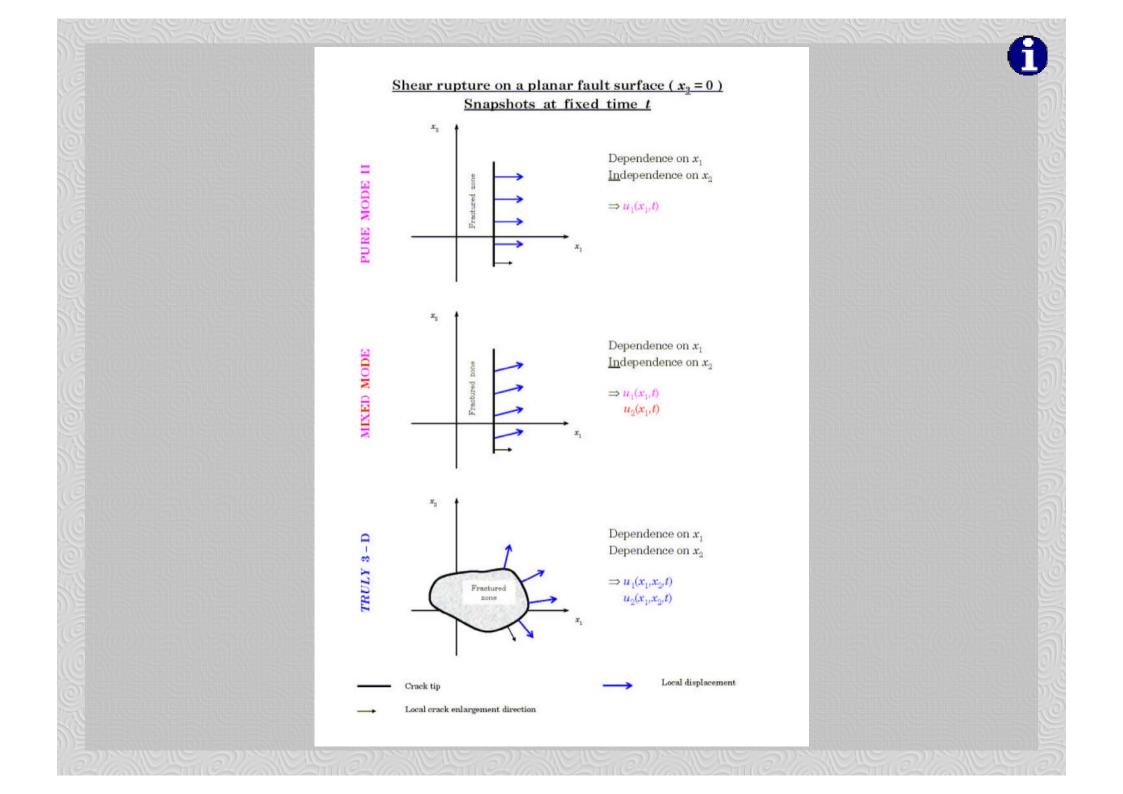
The vector  $U_i(\mathbf{x})$  can be reduced to a one – valued function if we produce in the body and starting from the dislocation curve a cut ( the **dislocation surface**), through which we assume an explicit discontinuity of **U**. Being  $\zeta$  the coordinate normal to the dislocation surface we can write:

$$\Delta U_i \equiv u_i = \lim_{\zeta \to 0^+} U_i(\mathbf{x}) - \lim_{\zeta \to 0^-} U_i(\mathbf{x}) = b_i$$

where  $\mathbf{b} = (b_1, b_2, b_3)$  is called Burgers' s vector.

In this framework the dislocation is described from a *microscopic / crystallographic point of view*.





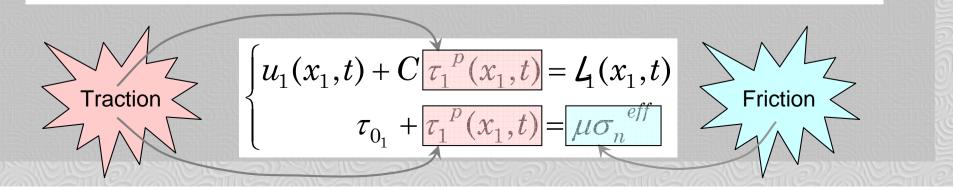
# Representation

#### **1.** INTEGRAL REPRESENTATION

Source integral rapresentation (*Betti*'s theorem, Integration in time (*Green – Volterra*'s relation), limit in fault surface, Lamb's problem):

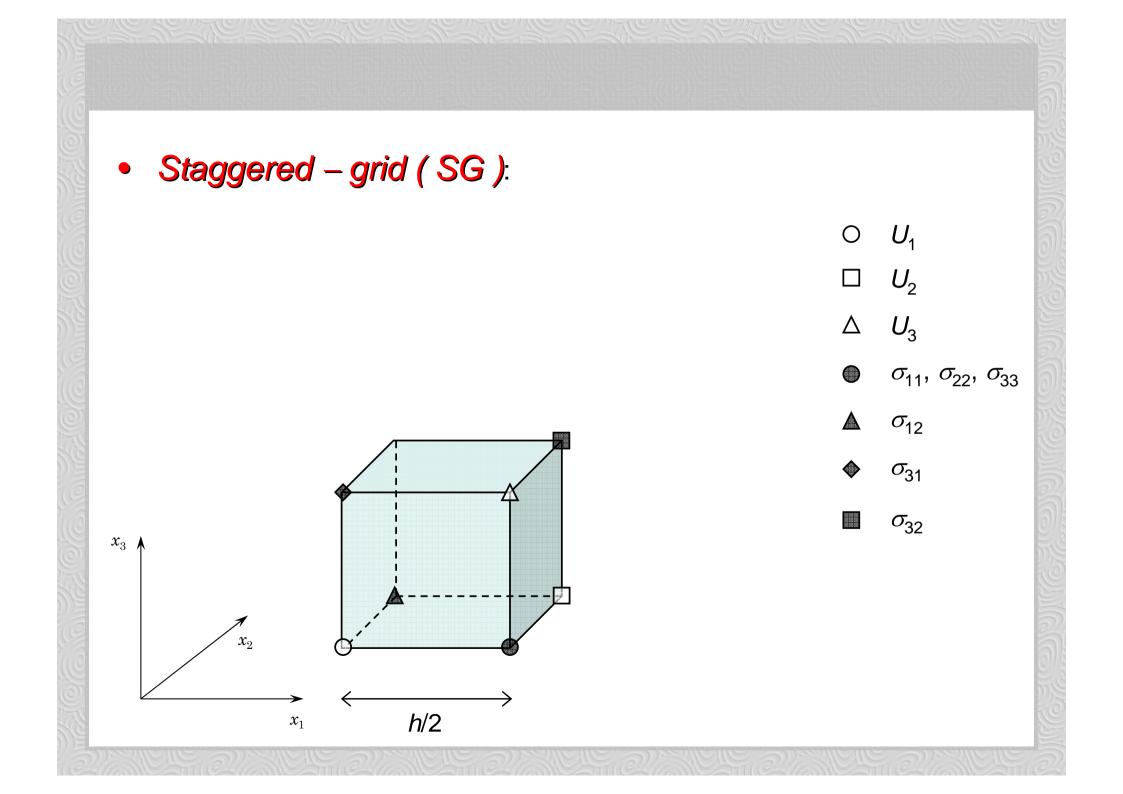
$$u_{n}(\mathbf{x},t) = \int_{-\infty}^{+\infty} dt' \int_{\mathcal{S}(t')} d\xi G_{n\alpha}(\mathbf{x}-\xi,t-t') \sigma_{\alpha\beta}^{p}(\xi,t') ; n=1,2,3; \alpha=1,2; \mathbf{x},\xi \in \mathbb{R}^{3}$$

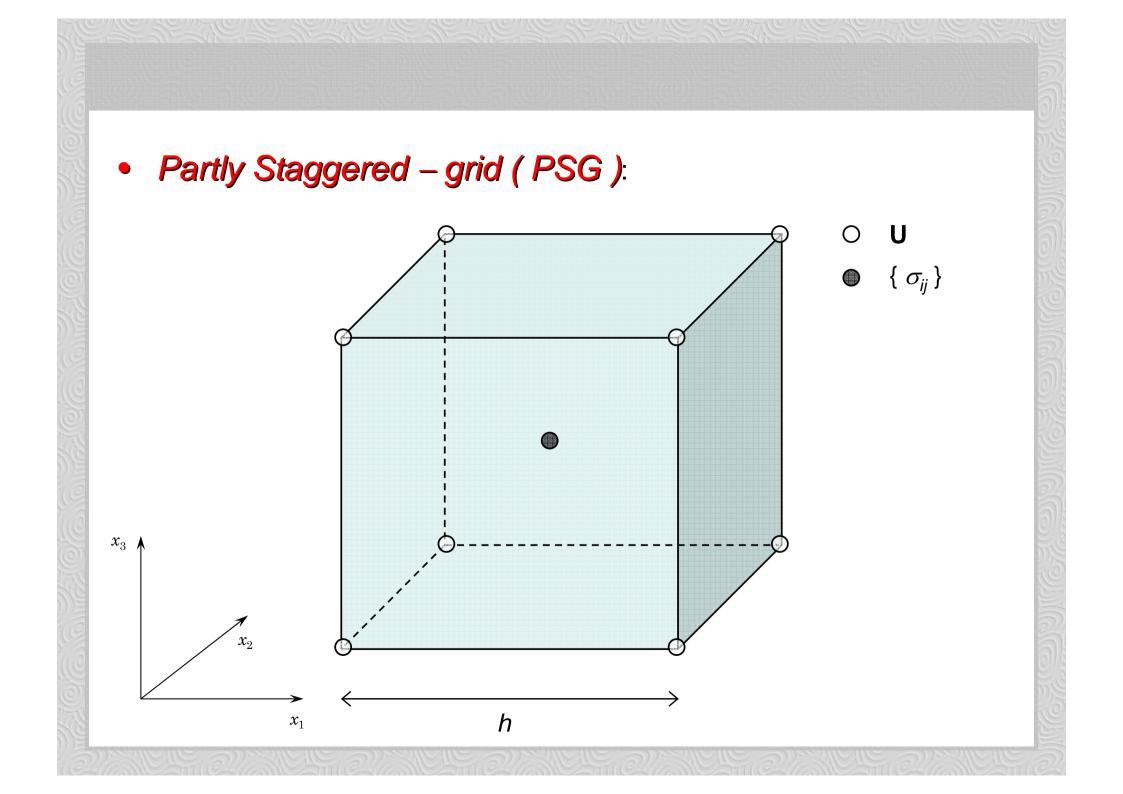
First neighbours decoupling ( in the case of a 2 - D, pure in - plane rupture ):

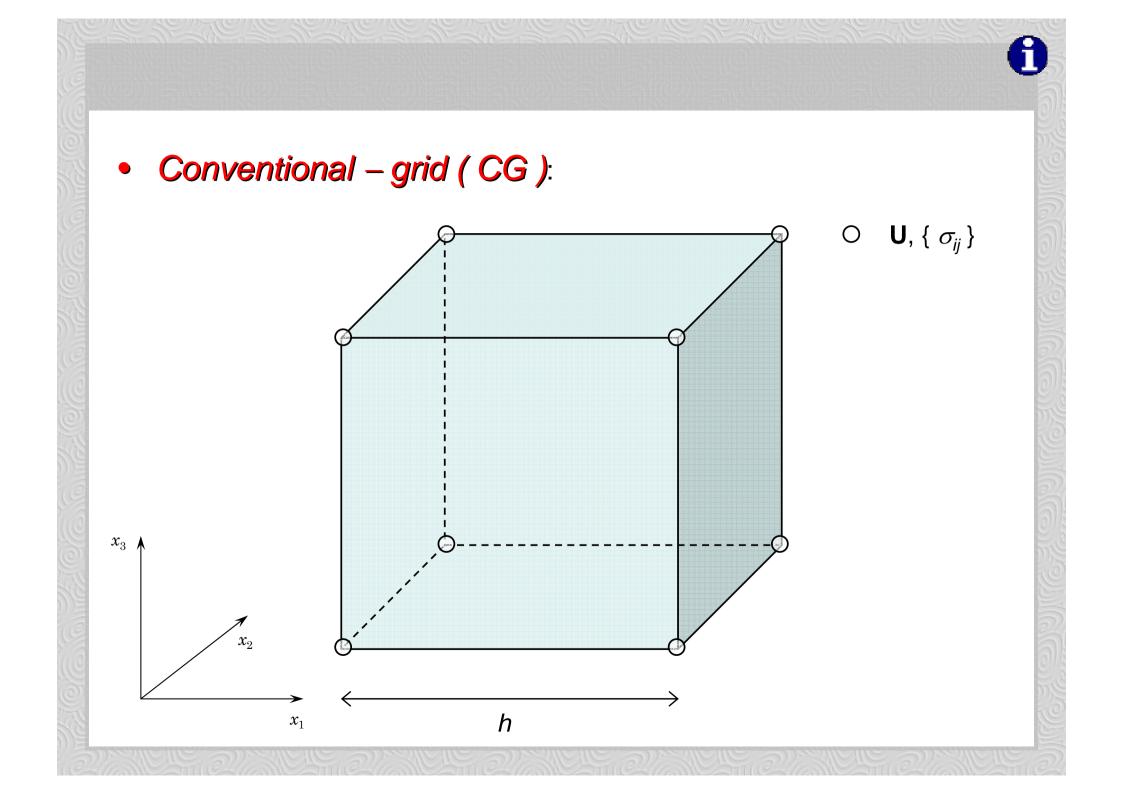


#### **2.** DISCRETIZATION OF EQUATIONS ( FE, FD APPROACHES )

Choice of the grid type







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- \* BOUNDARY:
  - mottoB
    - Fixed
    - Absorbing
  - ၂၀၀
    - Free surface
    - Topograph
    - Coasts
  - Lateral
    - Cyclic
    - Absorbing

• Let us consider a boundary perpendicular to the i – axis. Indeces i, j and k identify node location along  $x_1$ ,  $x_2$  and  $x_3$  axes, respectively. Apex m indicates the actual time level, while index l stands for vector component (l = 1, 2, 3).

#### • Fixed Boundary (FB):

$$U_{1\,jk_{l}}^{m} = 0, \quad \dot{U}_{1\,jk_{l}}^{m} = 0$$
  
 $U_{i_{end}jk_{l}}^{m} = 0, \quad \dot{U}_{i_{end}jk_{l}}^{m} = 0$ 

( Conditions  $\dot{U}^m_{1jk_l} = 0$  and  $\dot{U}^m_{i_{end}jk_l} = 0$  represent a Dirichlet boundary condition ).

#### • Absorbing Boundary (AB)

#### Left boundary:

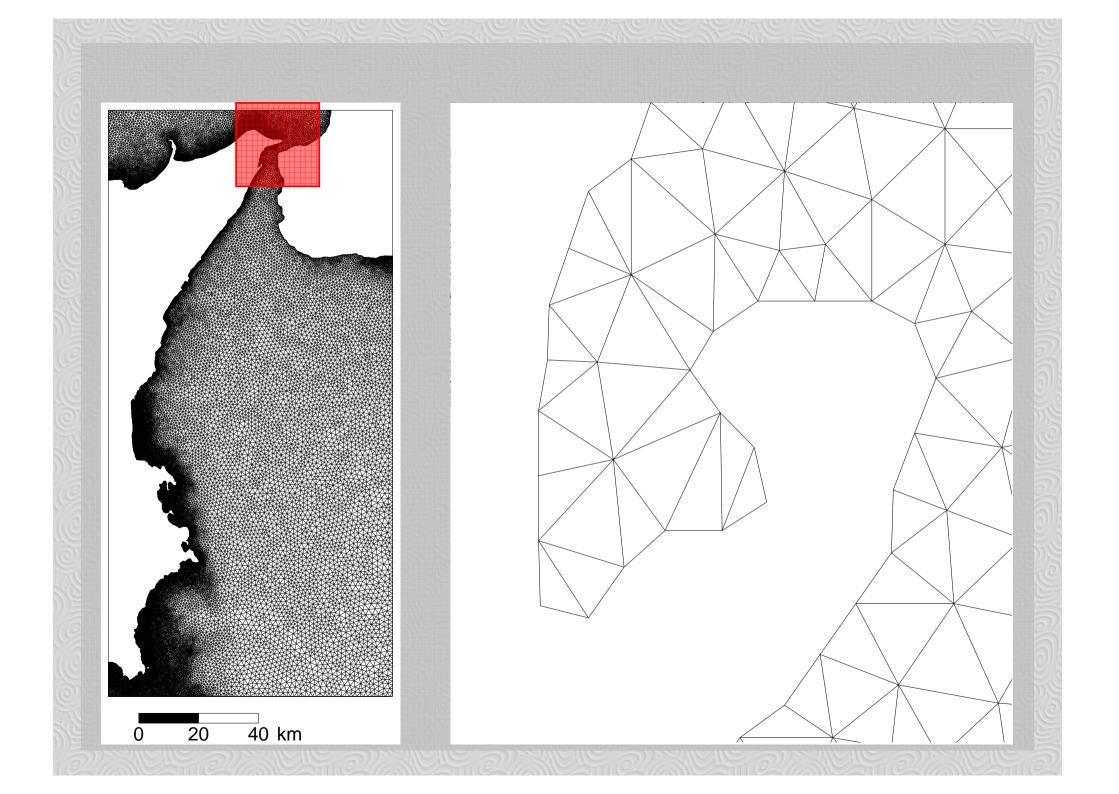
$$\begin{split} \dot{U}^{m}_{1\,jk_{l}} &= & A_{01}\dot{U}^{m}_{2\,jk_{l}} &+ A_{02}\dot{U}^{m}_{3\,jk_{l}} \\ &+ A_{10}\dot{U}^{m-1}_{1\,jk_{l}} &+ A_{11}\dot{U}^{m-1}_{2\,jk_{l}} &+ A_{12}\dot{U}^{m-1}_{3\,jk_{l}} \\ &+ A_{20}\dot{U}^{m-2}_{1\,jk_{l}} &+ A_{21}\dot{U}^{m-2}_{2\,jk_{l}} &+ A_{22}\dot{U}^{m-2}_{3\,jk} \end{split}$$

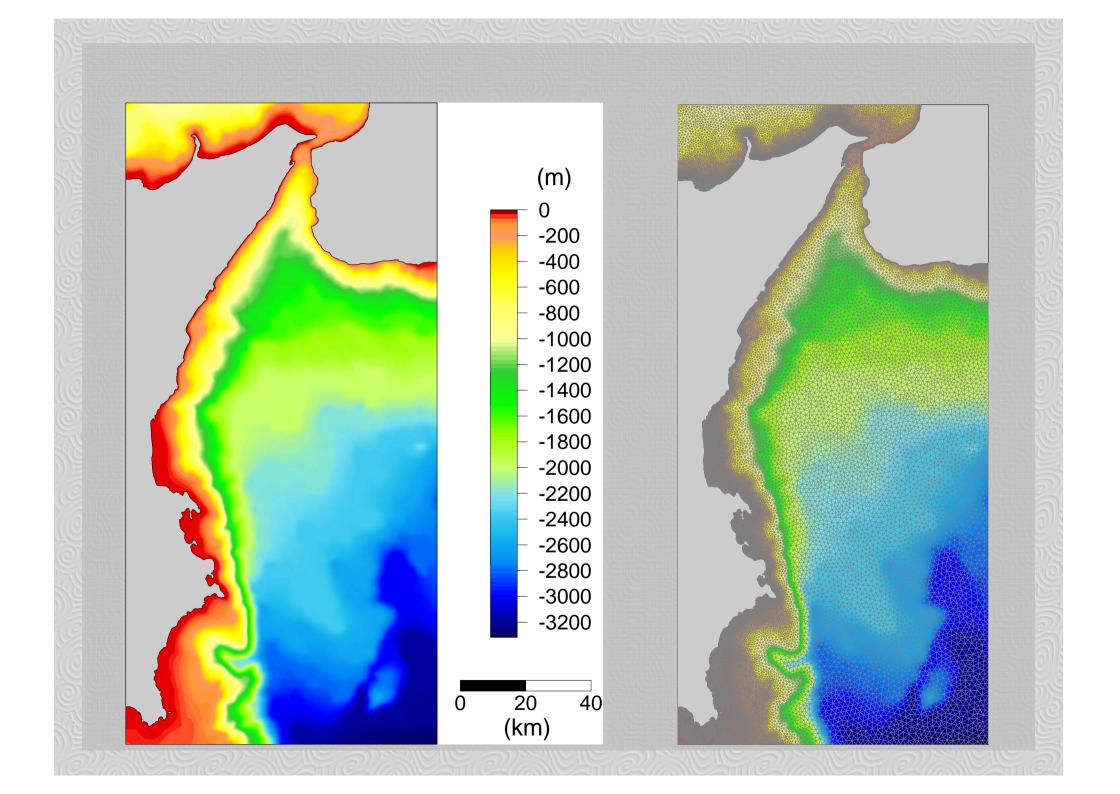
Right boundary:

$$\begin{split} \dot{U}_{i_{end}jk_{l}}^{m} &= & A_{01}\dot{U}_{i_{end}-1jk_{l}}^{m} + A_{02}\dot{U}_{i_{end}-2jk_{l}}^{m} \\ &+ A_{10}\dot{U}_{i_{end}jk_{l}}^{m-1} + A_{11}\dot{U}_{i_{end}-1jk_{l}}^{m-1} + A_{12}\dot{U}_{i_{end}-2jk_{l}}^{m-1} \\ &+ A_{20}\dot{U}_{i_{end}jk_{l}}^{m-2} + A_{21}\dot{U}_{i_{end}-1jk_{l}}^{m-2} + A_{22}\dot{U}_{i_{end}-2jk_{l}}^{m-2} \end{split}$$

In the previous compact representation of ABCs (that follows *Moczo*, 1998):

- the coefficients  $\{A_{pq}\}_{p,q=1,2,3}$  depend on the choice of ABC scheme (e. g. *Clayton and Engquist*, 1977; *Reynolds*, 1978; *Emerman and Stephen*, 1983; *Higdon*, 1991; *Peng and Toksöz*, 1994, 1995; *Liu and Archuleta*, 2000, ... );
- displacement components at actual time level *m* are derived by numerical integration from particle velocity components, after update;
- values in edges and in corners are derived from algebraic averaging of values of quantities belonging to walls;
- as it is a special boundary condition, there is no need to consider any rheology in the updated point.





Number of nodes	30264
Number of elements	57733
Type of building block	Triangle
Minimum node distance	200 m
Maximum node distance	2000 m
References	Armigliato A., Tinti S., (2005), EGU General Assebly;
	Tinti S., Armigliato A., Bortolucci E. (2001), <i>J. Seismol.</i> , <b>5</b> , 41-61.

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# Fault Boundary Conditions

#### **1.** *ТҮРЕ*

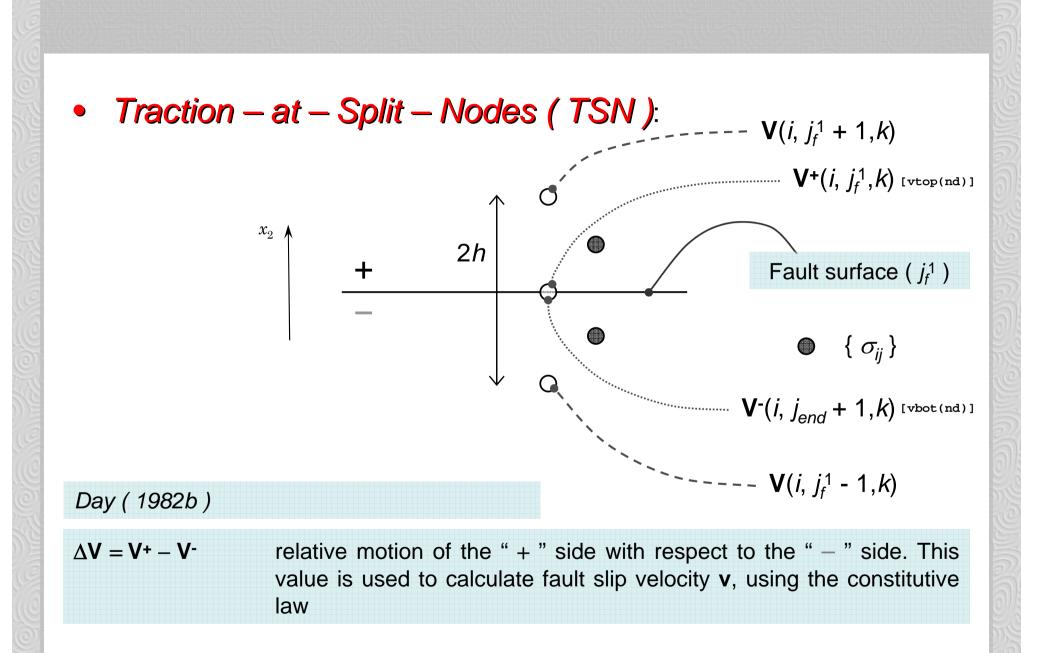
- Traction at Split Nodes (**TSN**): in **2 D** by Andrews (1973); in **3** – **D** by Day (1977), Archuleta and Day (1980), Day (1982a, 1982b), Andrews (1999), Bizzarri (2003), Bizzarri and Cocco (2005), Day et al. (2005)
- Stress Glut ( **SG** ): Backus and Mulchay ( 1976 ), Andrews ( 1976 )

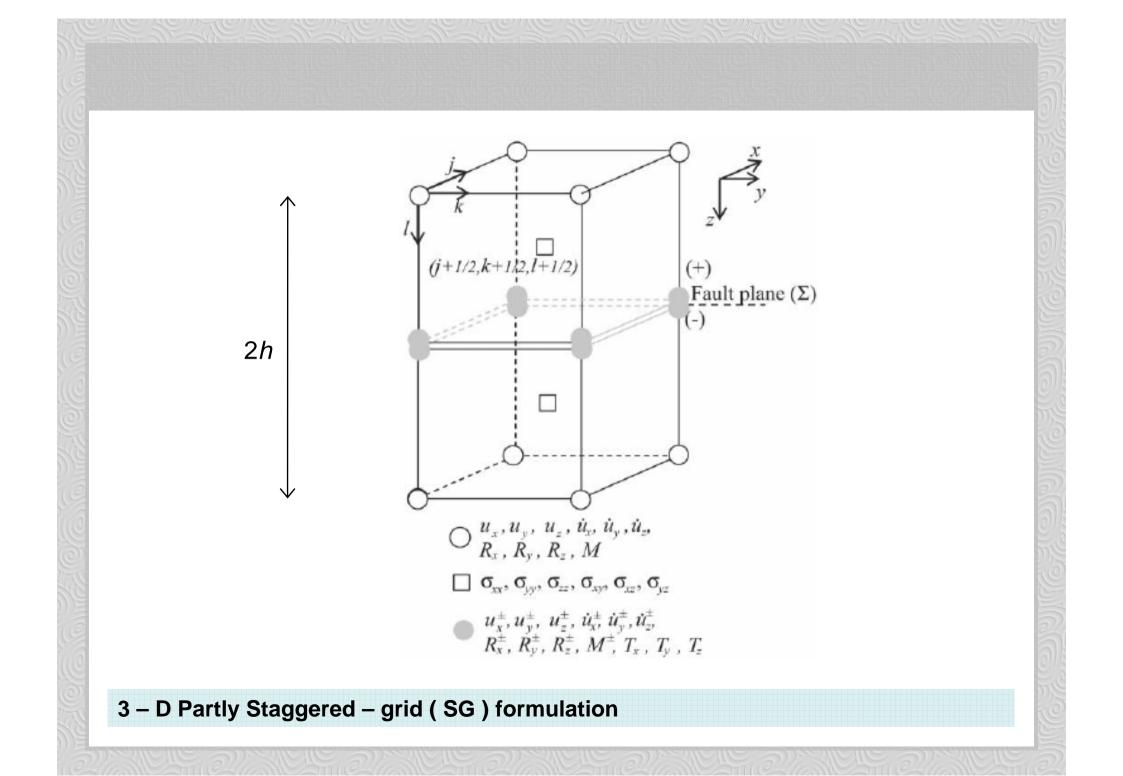
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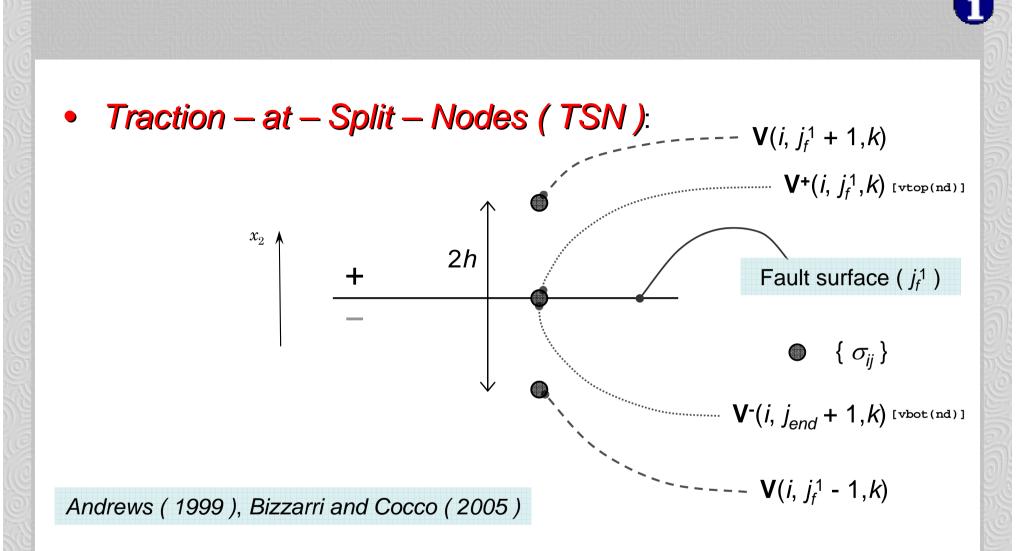
- Thin zone (**TnZ**): *Virieux and Wadariaga (1932)*
- Thick zone (TkZ): Madariaga et al. (1998)

#### 2. CONSTITUTIVE LAW

- Accounts for fault rheology
- and different physical phenomena occurring during the rupture process

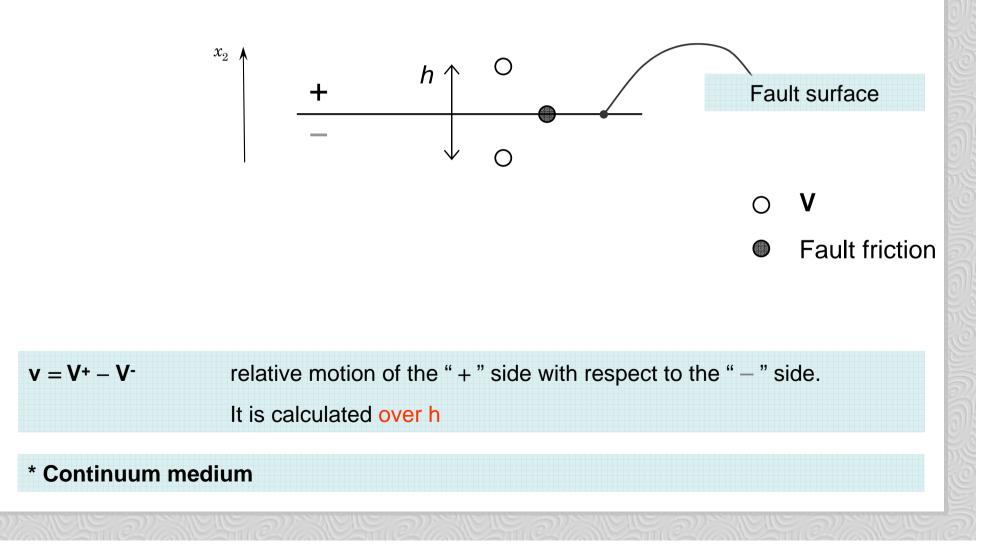


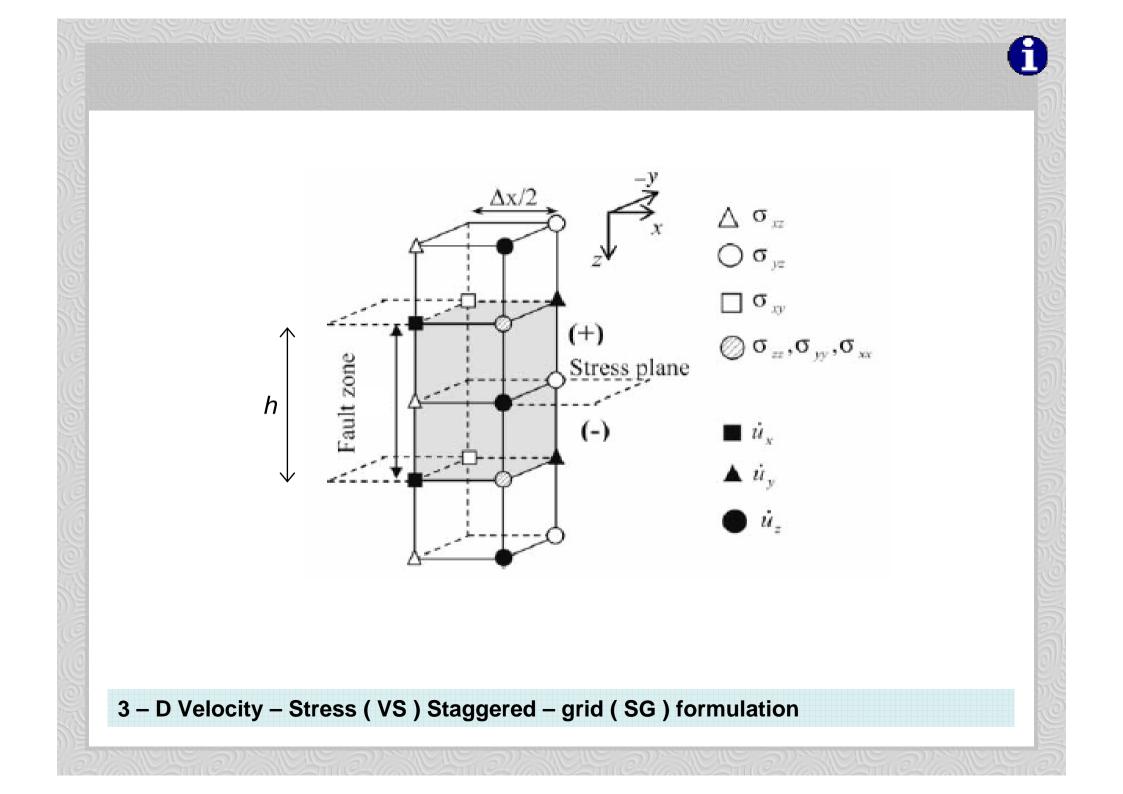




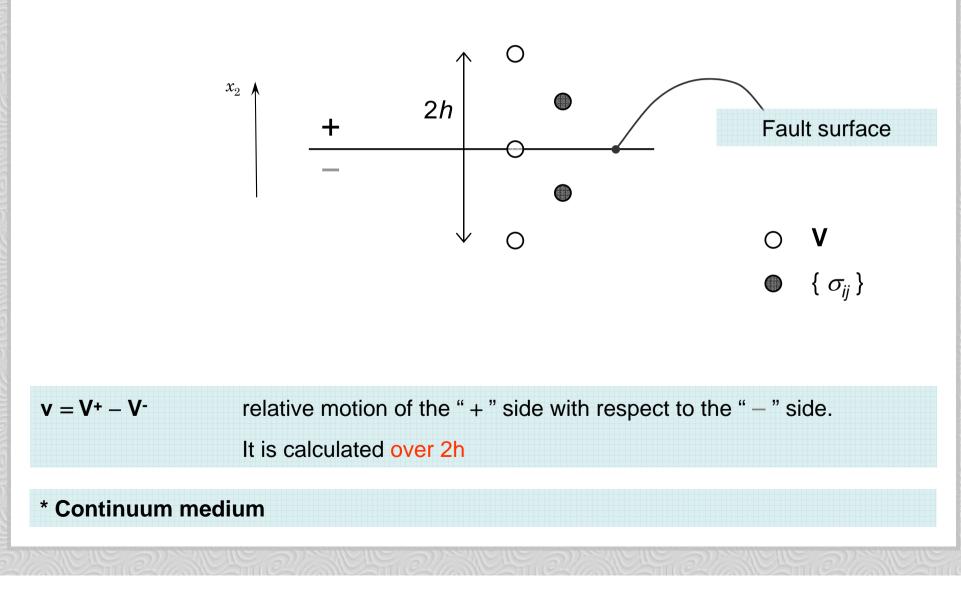
\* **Discontinuum medium** ( continuum mechanics equations of motion are applied to each half - space individually; the fault is an explicit discontinuity in displacement )

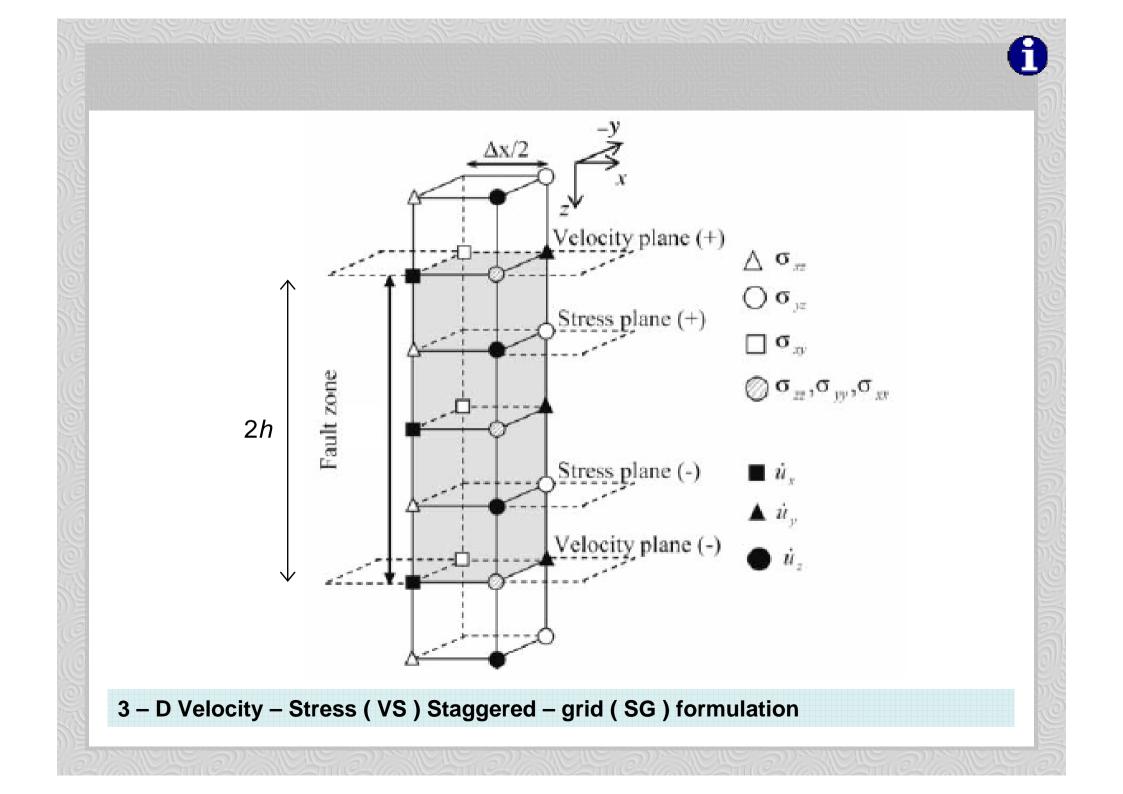
• Thin -zone(TnZ):





• Thick – zone (TkZ)





# Auxiliary Conditions

\* COLLINEARITY BETWEEN FAULT SHEAR TRACTION AND FAULT SLIP VELOCITY:

Τ // ν

(i.e. 
$$\hat{\mathbf{T}} = \frac{\mathbf{v}}{\|\mathbf{v}\|}$$
).

